Investigation of the contribution of topographic effects on regional geoid modeling within the Geomed2 project

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Introduction

With nowadays high resolution Global Geopotential Models (GGMs), traditional geoid approximation with the Remove-Compute-Restore (RCR) procedure needs revisiting both in terms of the numerical and the methodological steps followed. This becomes evident if one considers the high degree of expansion of the latest combined GGMs, corresponding to frequencies as short as 9.2 km for a maximum degree of expansion equal to 2160, and the limited spatial resolution of bathymetry models, being 30 arcsec (~1 km) for the SRTM15_PLUS model. In the above schema, if a high-resolution GGM is used to reduce the available surface gravity data, then the bathymetry model used for the terrain reduction should be able to represent adequately the spatial frequencies between 10 and 1 km, so that the corresponding gravity signal presents in the reduced gravity anomalies can be removed. This step is quite significant since depending on the correctness, resolution and accuracy of the bathymetry model, the resulting residual gravity anomalies can be either useful for geoid determination, or can introduce biases, noise exaggeration and aliasing in the predicted gravimetric geoid model. Within the GEOMED 2 project, which aims at a high resolution and accuracy geoid determination for the entire Mediterranean basin, in situ gravity anomalies are used within the RCR procedure employing DIR-R5 and EIGEN6c4 as reference geopotential models and various bathymetry/topography models (DTU10, EMODNET and SRTM15_PLUS) in order to investigate the influence of the used terrain representation on the gravity and geoid signal. The used Digital Terrain and Bathymetry Models (DTBMs) are of various resolutions, smoothness, and accuracy, while some of them are directly derived from satellite data and some of them from in situ echo sounding observations. In this work the Residual Terrain Model (RTM) reduction is used to model the contribution of topography/bathymetry to gravity and the geoid, while the contribution of each model and its appropriateness for use i

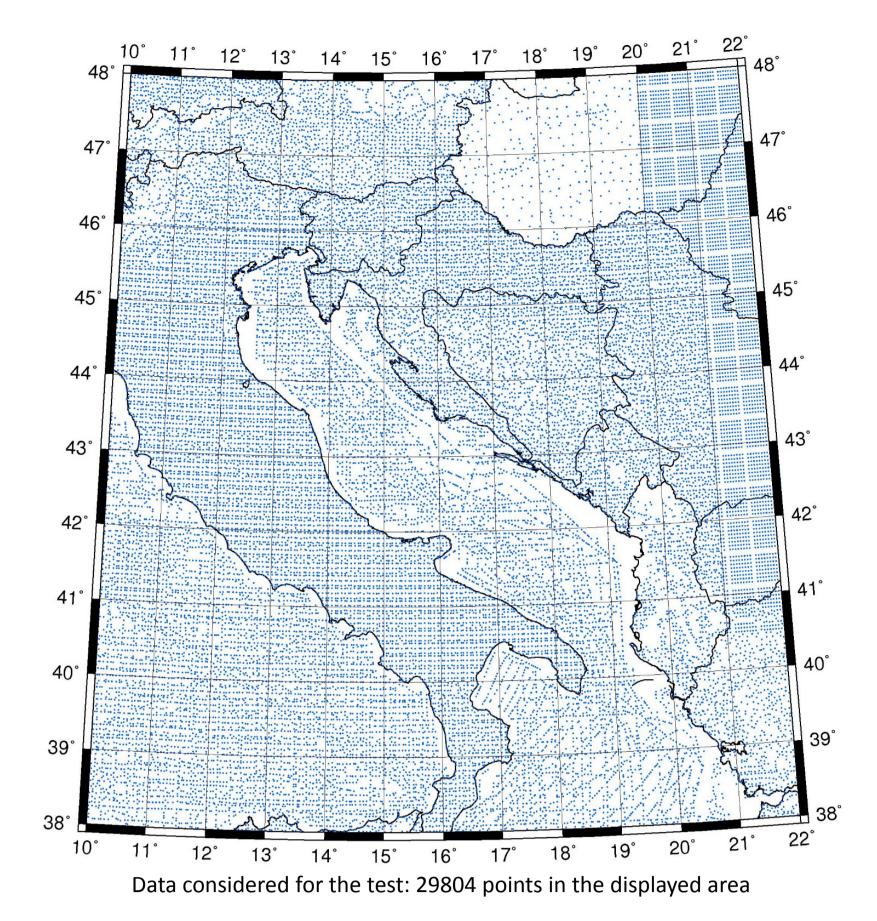
Digital Terrain and Bathymetry Models (DTBMs)

Over land the SRTM3 has been used. Over sea the following models have been considered.

DTU10 (1'x1'): Bathymetry derived from satellite ERS-1 data, mapped with a resolution of 1 minute by 1 minute corresponding to 2 minute by 2 minute resolution at Equator (Andersen et al., 2008).

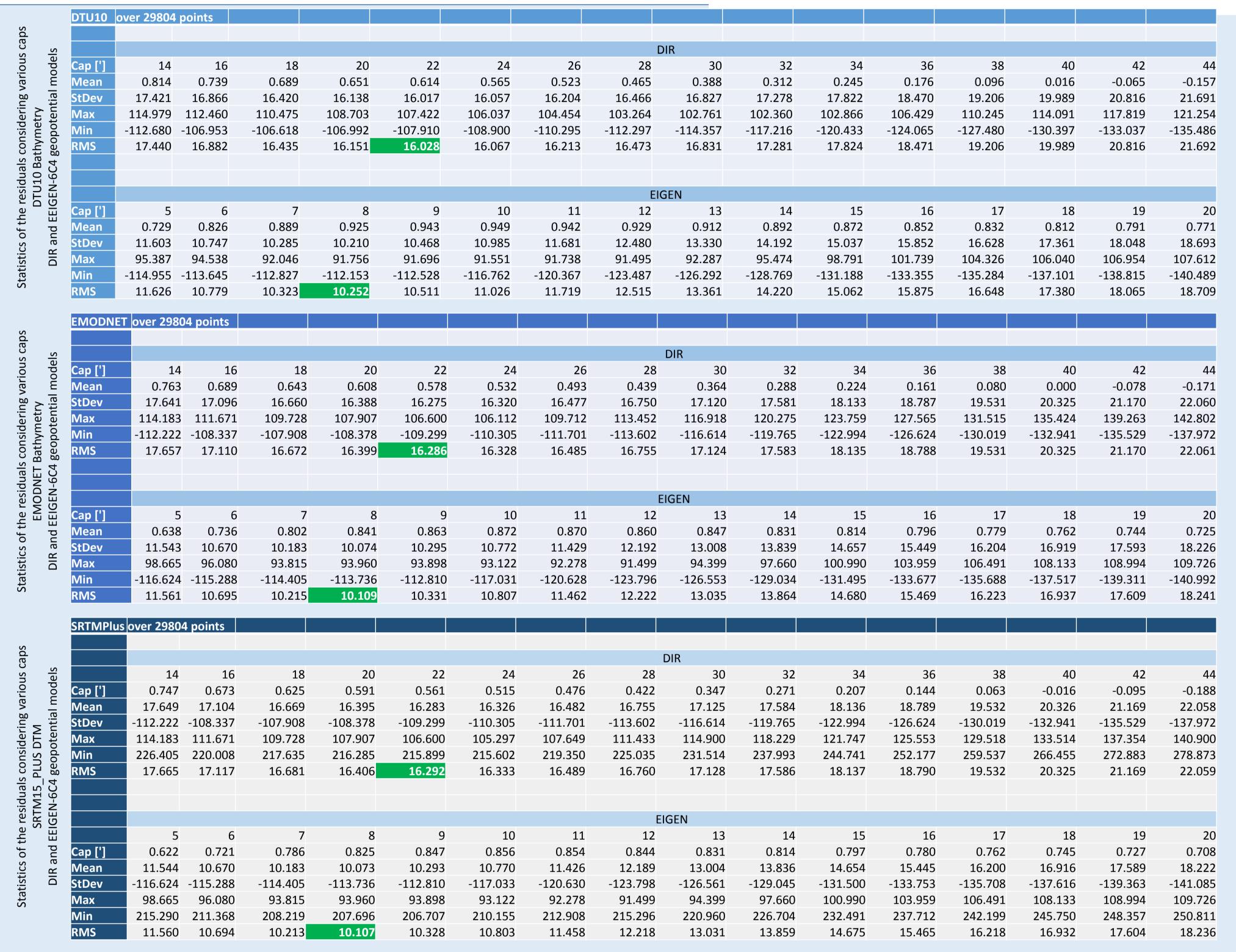
EMODNET (7.5"x7.5"): The EMODnet-Bathymetry portal (http://www.emodnet-bathymetry.eu) is being developed in the framework of the European Marine Observation and Data Network (EMODnet) as initiated by the European Commission. There are 4 types of organisations (Hydrographic Offices, Authorities, Research institutes, Industry) that perform bathymetric surveys, thereby partly overlapping and mostly complementing their geographical coverages. Data are collected at different frequencies and even date back to previous centuries. The partners combine expertise and experiences of collecting, processing, and managing of bathymetric data together with expertise in distributed data infrastructure development and operation and providing OGC services (WMS, WFS, and WCS) for viewing and distribution.

SRTM15_PLUS (15"x15"): Land elevations are based on the best available data from SRTM, ASTER digital elevation models while the ice topography of Greenland and Antarctica is based on CryoSat-2 and IceSat. Ocean bathymetry is based on bathymetric predictions from the latest global gravity model from CryoSat-2 and Jason-1 along with 494 million edited depth soundings at 15 arcsec resolution. Bathymetry of the Arctic seafloor is based on the IBCAO grid with improved resolution in areas of multibeam coverage. This SRTM15_PLUS provides the foundational bathymetry layer for Google Earth and is freely available at the topex.ucsd.edu ftp site (Olson and Sandwell, 2016).

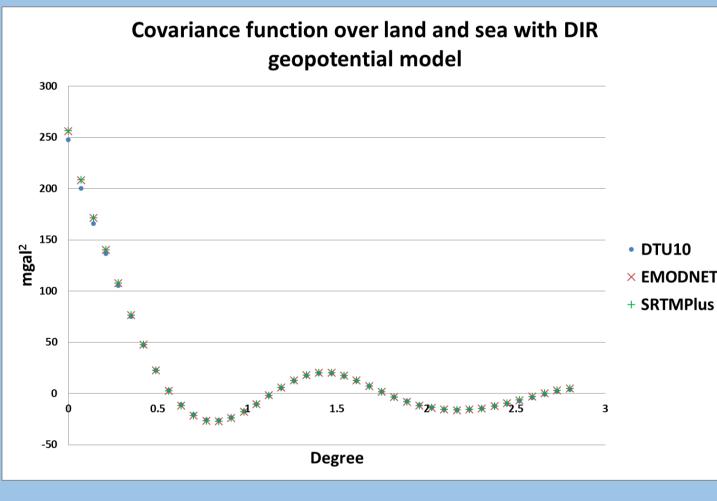


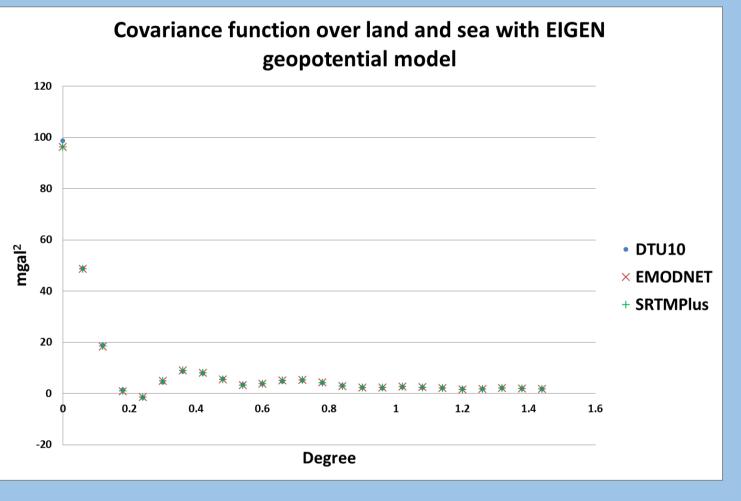
Statistics of the residuals Δg after the Remove phase. Different Geopotential Models (DIR up to d/o 230 and EIGEN-6c4 up to d/o 1000) have been used and Digital Terrain Models including different bathymetries have been considered. Various caps for getting the coarse DTM to be used for the Residual Terrain Correction component have been considered in the computation, to find the one providing the best reduction (highlighted in green).

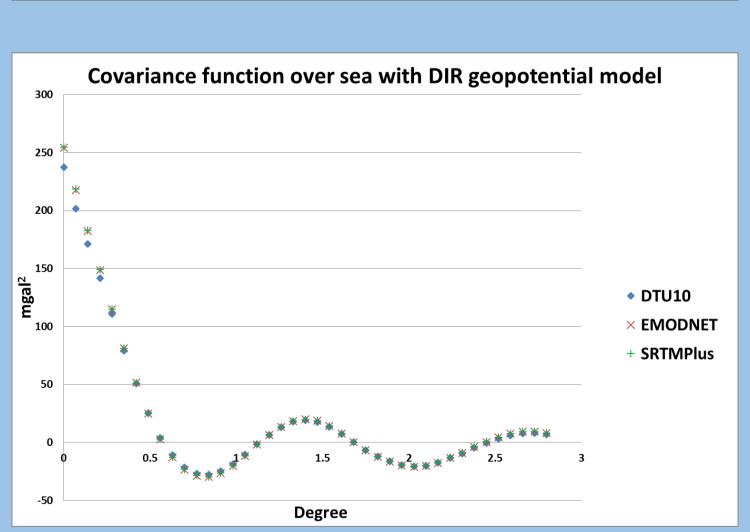
The best reduction is obtained with a 22' cap for DIR geopotential model and with a 8' cap for EIGEN-6c4. The results are equivalent for the three considered bathymetries.

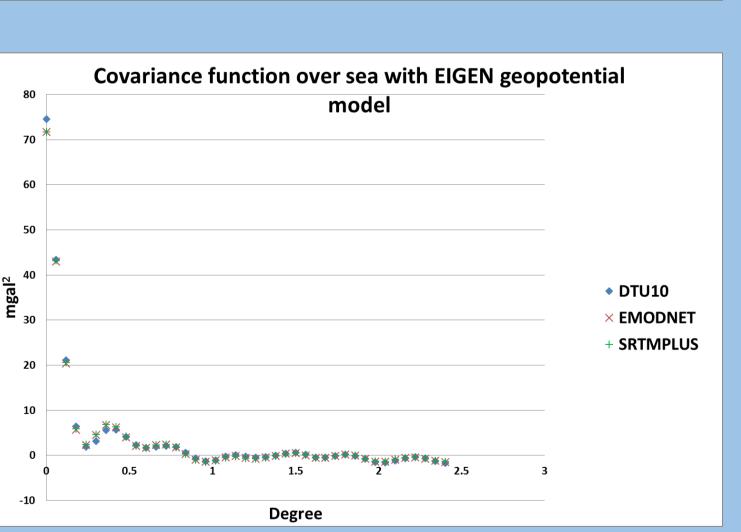


Covariance function over the whole area (land and sea) or over sea only, considering different geopotential models (DIR and EIGEN-6c4) and using DTMs based on different bathymetries. It is possible to observe that the behaviour of the covariance function is less regular than expected. The effect does not vary with respect to the considered bathymetry.



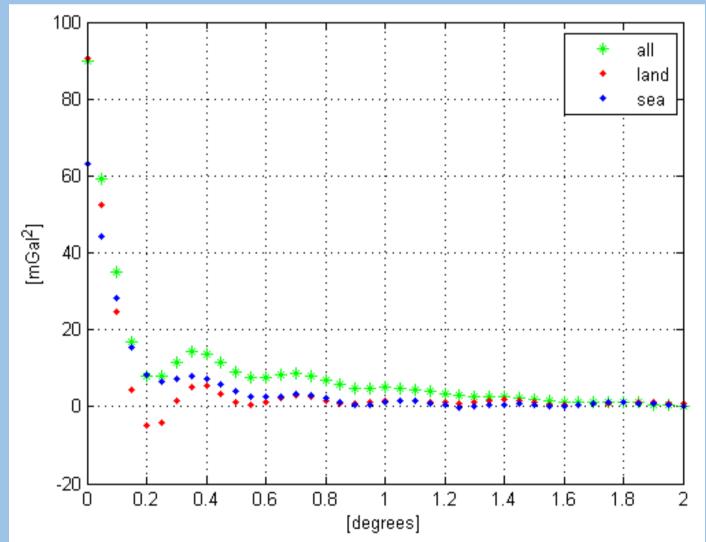






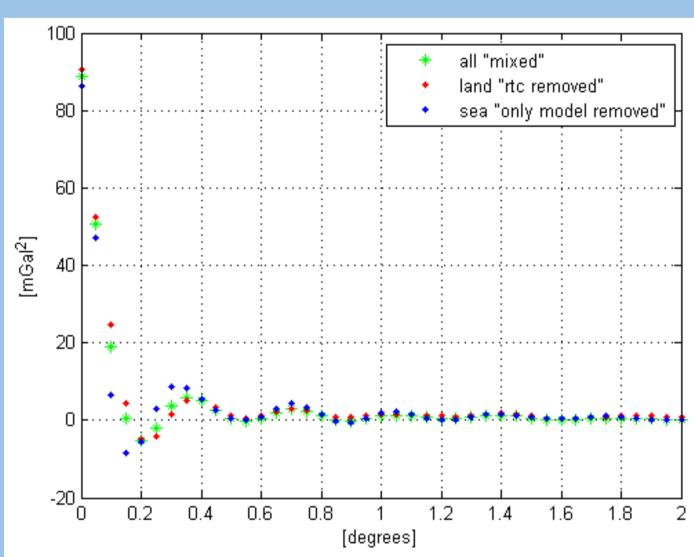
the whole area (land and sea) or over land only or sea only, considering EIEGEN-6c4 geopotential model and the DTM based on EMODNET bathymetry. It is possible to observe that the covariance function over land only has a more regular behaviour.

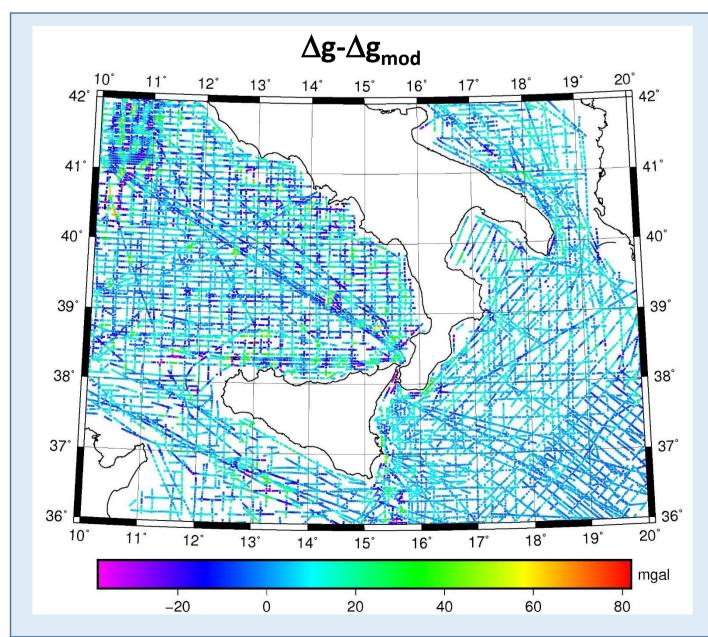
Covariance function over



Covariance function over the whole area (land and sea) or over land only or sea only, considering EIEGEN-6c4 geopotential model and the DTM based on EMODNET bathymetry.

bathymetry. In this case the data over land have been completely reduced ($\Delta g - \Delta g_{model} - \Delta g_{RTC}$), while the data over sea have been only partially reduced ($\Delta g - \Delta g_{model}$). It is possible to observe that the behaviour of the covariance function has improved, implying that the matter is to be ascribed to the RTC component over





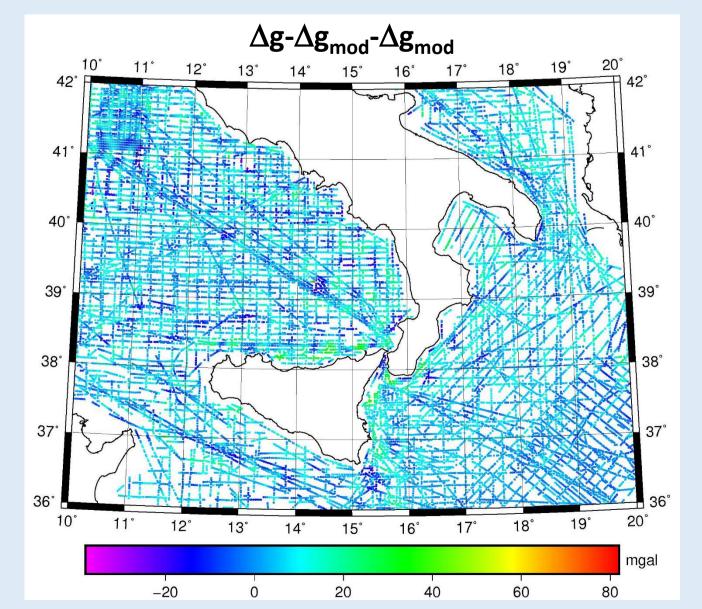


Image and statistics of the residuals partially reduced (Δg - Δg_{model}) and completely reduced (Δg - Δg_{model} - Δg_{RTC}), considering EIEGEN-6c4 geopotential model and the DTM based on EMODNET

bathymetry.

After removing also the RTC component a high frequency signal is still present in the residuals, suggesting that some of the gravity contribution of the terrain cannot be modelled with the available information.

	Δg - Δg_model	Δg – Δg_{model} - Δg_{RTC}
# Values	22966	22966
Mean	3.160	3.775
St Dev	9.794	7.280
Min	-53.942	-38.399
Max	84.816	74.946

Conclusions

In view of the new computation of the Mediterranean geoid with the Remove-Compute-Restore technique, a test has been performed to investigate the performance of the reduction of gravity data considering different geopotential models and DTMs based on different bathymetries. The observation of the reduced data, of their statistics and of the covariance functions have shown that there are possible mismatches between marine data and bathymetry.