



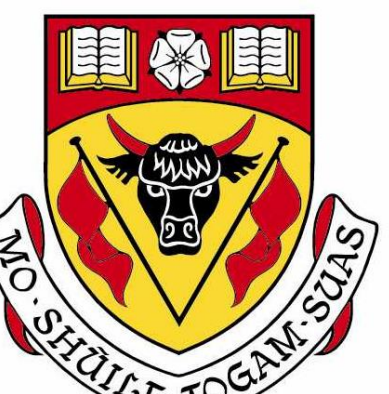
Combination of GOCE SGG data and surface gravity anomalies for local/regional geoid determination

R.S. Grebeticharsky¹, G.S. Vergos², D. Piretzidis³, I.N. Tziavos²

¹RSG Consult Ltd., Bulgaria, rgrebenitcharsky@gmail.com

²Department of Geodesy and Surveying, Aristotle University of Thessaloniki, Greece, vergos@topo.auth.gr

³Department of Geomatics Engineering, University of Calgary, Canada



UNIVERSITY OF CALGARY

Introduction

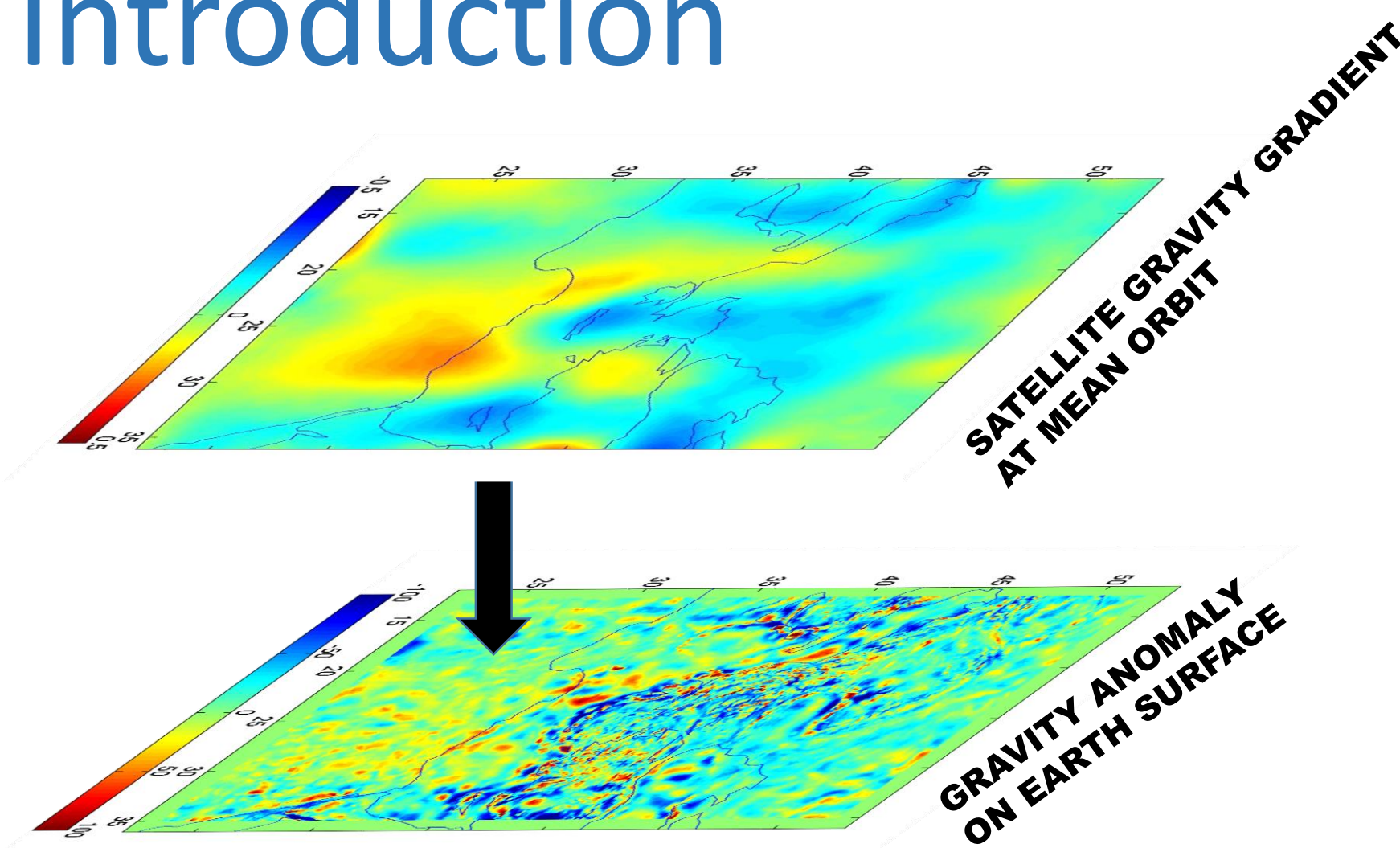


Figure 1: MAIN TASK DESCRIPTION

MAIN TASK: Combination of SGGs at mean orbit and ground gravity anomaly data on the Earth's surface through GOCE gradient downward continuation for local/regional geoid modeling:

- Using a Monte Carlo Method (Simulated Annealing) for solution of inverse problems
- Application of Quasi-random generator (QG)
- Validation of geoid model by GPS&GNSS/Leveling

Background

MOTIVATION: TO COMBINE SGG AT MEAN ORBIT WITH GRAVITY ANOMALY DATA ON THE EARTH SURFACE FOR REGIONAL GEOID MODELLING

REGIONAL GEOID MODELING: for performing calibration of SGGs with external gravity information on a grid to use all available data at highest possible resolution

- HOW? By application of a Monte-Carlo method (Simulated Annealing). Simulated annealing (SA) is a probabilistic method proposed by Kirkpatrick, Gelett and Vecchi (1983) for finding the global minimum of a cost function that may possess several local minima. It works by emulating the physical process whereby a solid is slowly cooled so that when eventually its structure is 'frozen', this happens at a minimum energy configuration.
- WHY? SA allows solving inverse problems like downward continuation of SGGs from mean orbit to the Earth's surface using an iterative Monte Carlo procedure based on quasi-random generator.
- WHAT TO DO? 1) Quasi-random generation of gravity anomalies on earth surface in the form of grid; 2) Upward continuation to mean orbit by MIMOS – Forward step; 3) Comparison of SGG observations with upward SGGs and 'freeze' those satisfying SA criteria; 4) Repetition of 1-3 till all generated SGGs meet the SA criteria.

FORWARD STEP : MULTIPLE INPUT OUTPUT SYSTEMS (MIMOS) FOR UPWARD CONTINUATION (TO THE MEAN ORBIT) (SIDERIS 1996 - MODIFIED)

$$FFT\{T_{ij}^{MO}\} = FFT\{K_{ij}^{MO,ES}(LT^{ES})\} \cdot FFT\{LT^{ES}\} \quad i, j = x, y, z$$

where : FFT - Fast Fourier Transform;

T_{ij}^{MO} - the gravity gradients at mean orbit along the axis i, j

LT^{ES} - Functional of disturbing gravity potential (gravity anomaly) on Earth surface

$K_{ij}^{MO,ES}$ - Kernel for transformation of the functional of disturbing potential LT^{ES} on Earth Surface to the gravity gradient T_{ij}^{MO} at mean orbit (Eshagh, 2010)

where: ES – Earth's surface; MO – mean satellite orbit at 260 km
{*} – element by element multiplication

Experiment description and results

- Application area: $22 < \text{latitude} < 53$; $12 < \text{longitude} < 36$
- Applied on: 2D regional $6' \times 6'$ grid for T_{zz} (directly observed) and $T_{zz,lp}$ derived from Laplace's equation with directly observed T_{xx} and T_{yy} :
 $T_{zz,lp} = -T_{xx} - T_{yy}$
- Gridding – standard procedure with Delaunay triangulation
- Time interval covered: 01/11/2009 – 30/04/2011 (16 months) with 421 days of SGG data
- Data type and source: EGG_NOM_2, GOCE HLPF, ESA/ESRIN

- SGG description: normal potential gradients are evaluated at the LNOF
- transformed to the GRF. Disturbing potential gradients are computed and then filtered with an FIR band-pass filter (5-100 mHz) with a Hamming window for spectral leakage. The filtered T_{ij} are then reduced to a mean orbit.
- Reference Gravity Field Model: GOCE DIR R4
- Ground data: EGM 2008 Gravity Anomaly Model generated on Earth's surface
- Type of model SGG data: DIR R4 GPM generated directly at mean orbital altitude
- Number of separate runs: 31
- Number of iterations necessary to reach SA criteria per every run: 500
- Validation of determined geoid models using GPS & GNSS/Leveling bench marks
- Number of GPS/Leveling benchmarks over land only: 1542

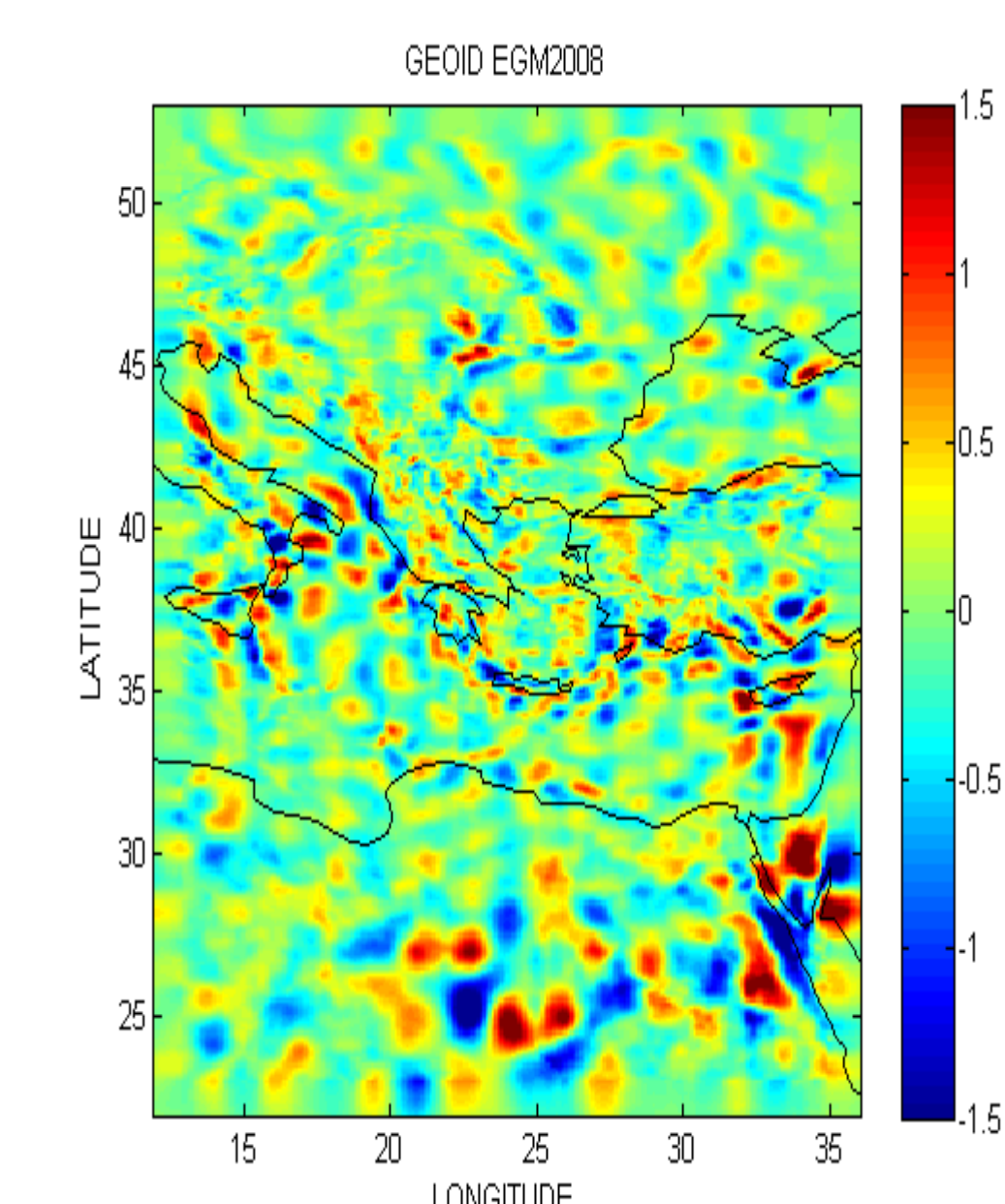


Figure 2: EGM 2008 geoid [m]

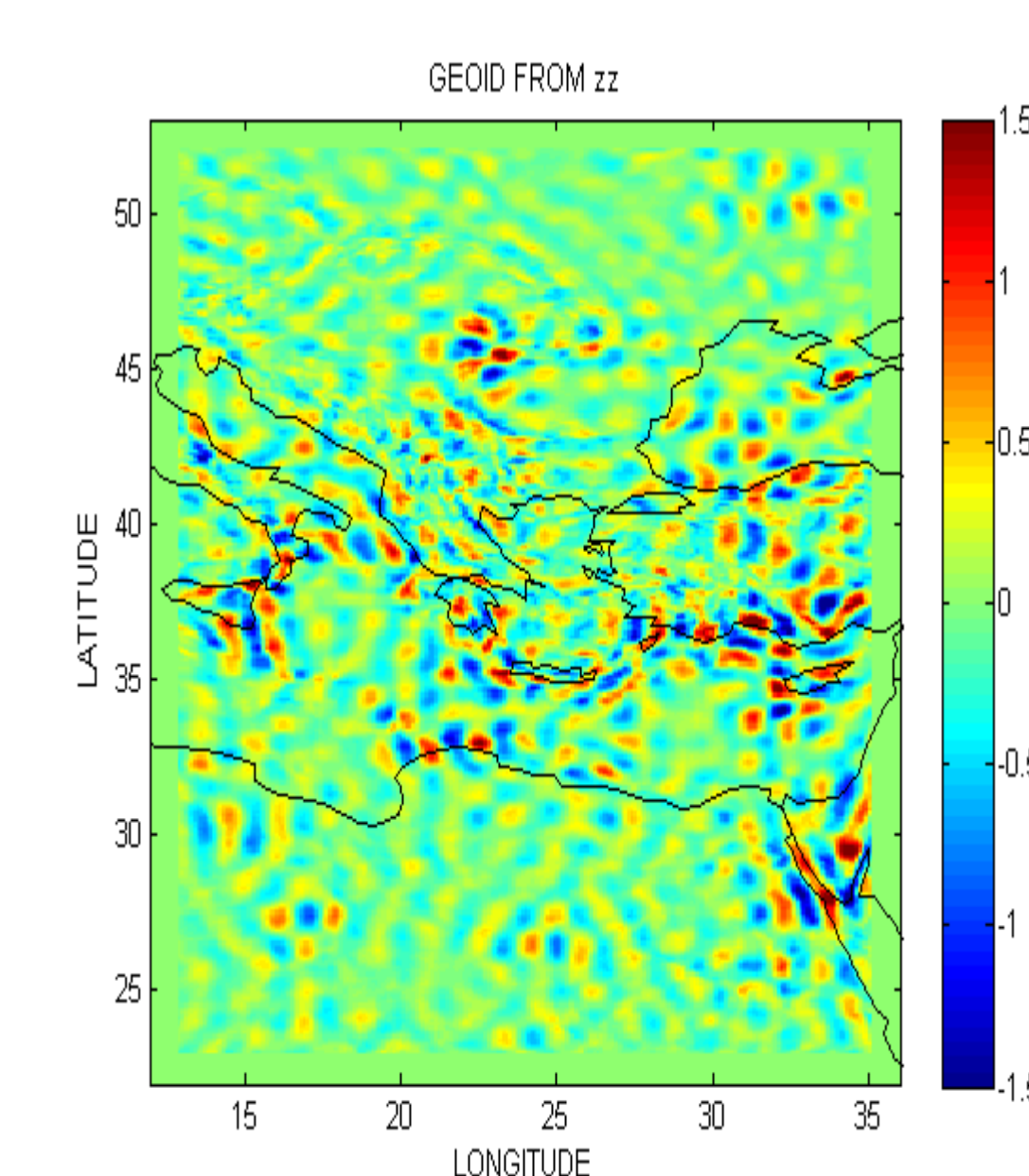


Figure 3: Geoid determined using T_{zz} [m]

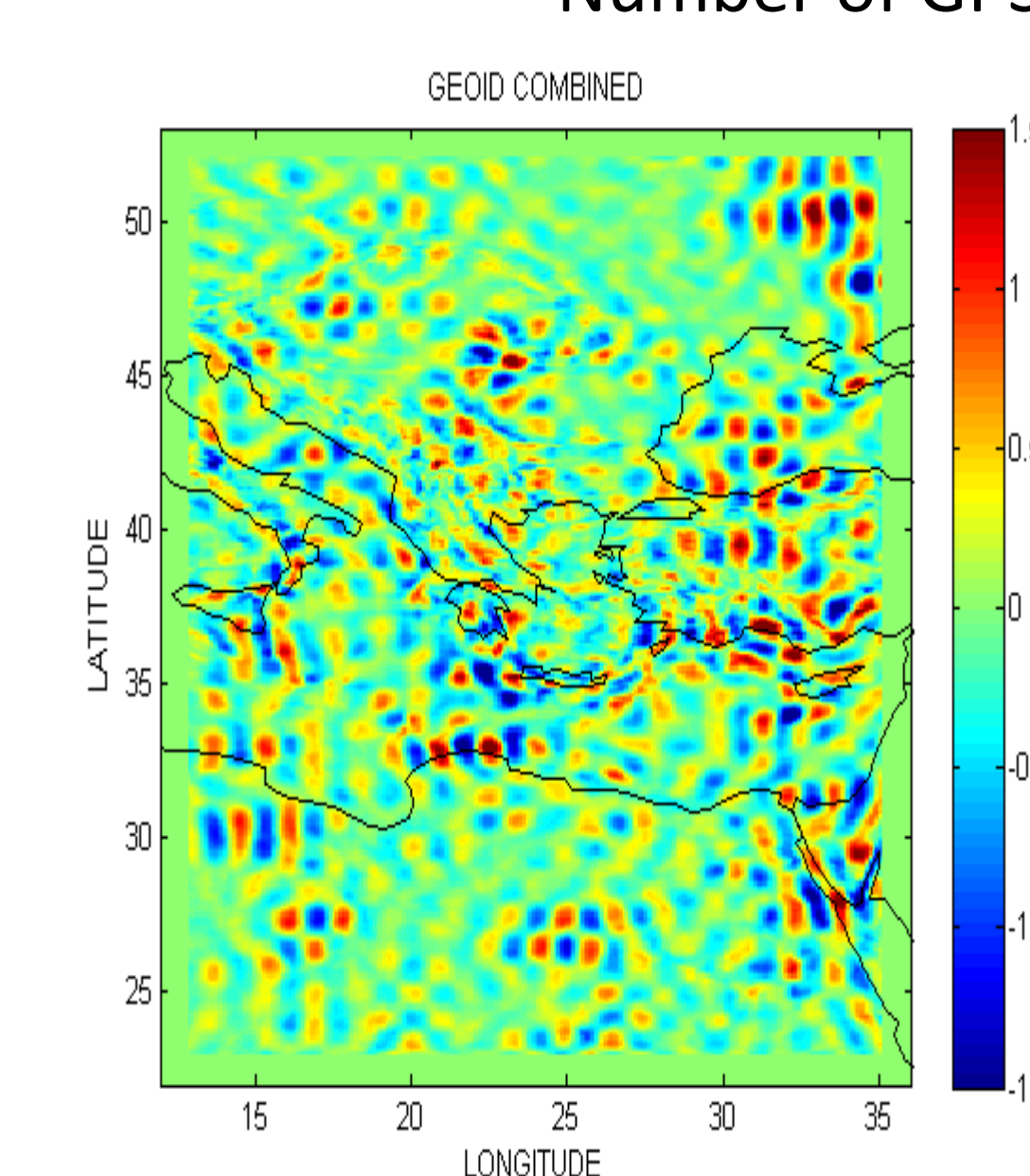


Figure 4: Geoid determined using $T_{zz,lp}$ [m]

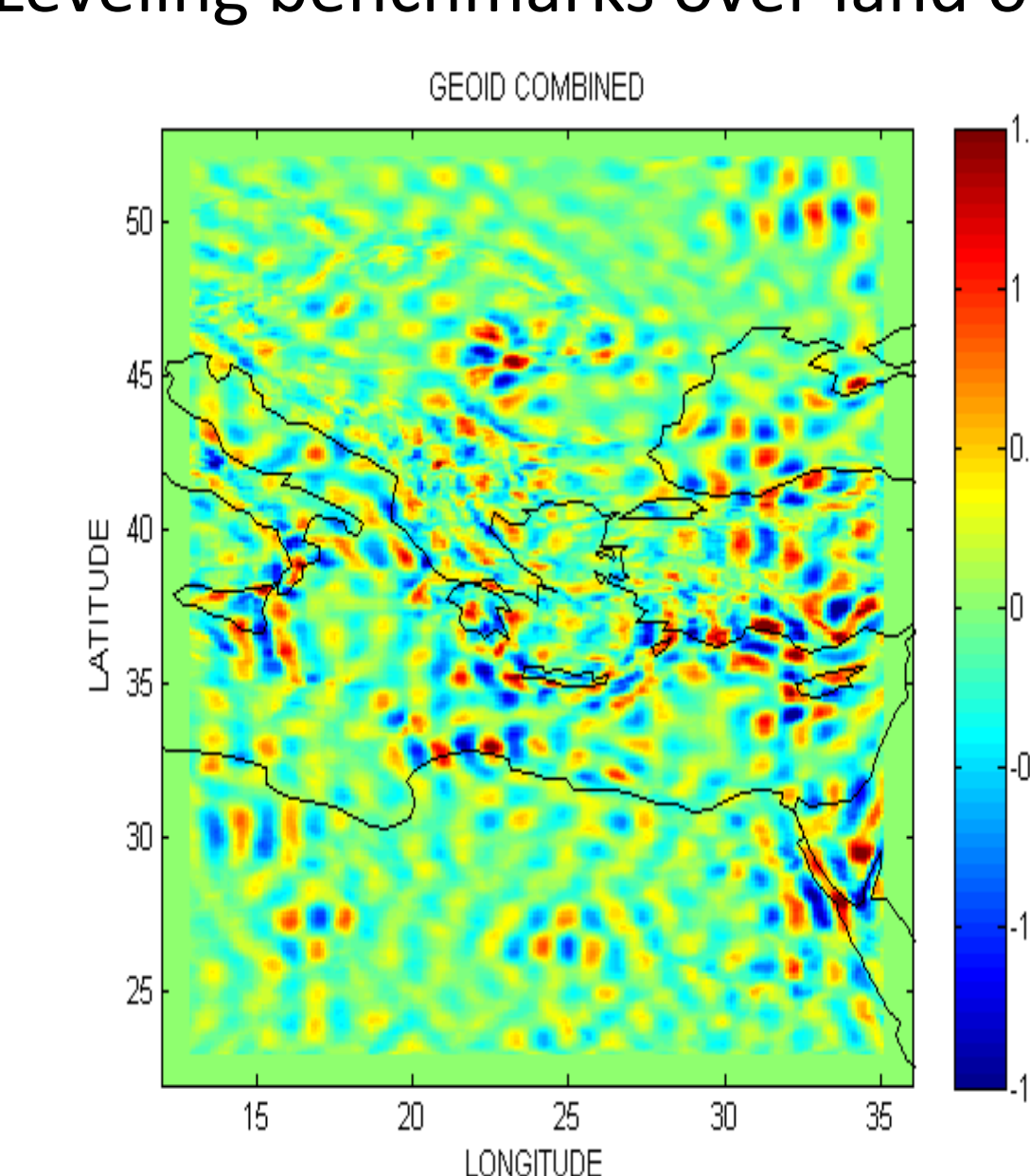


Figure 5: Combined Geoid (T_{zz} & $T_{zz,lp}$) [m]

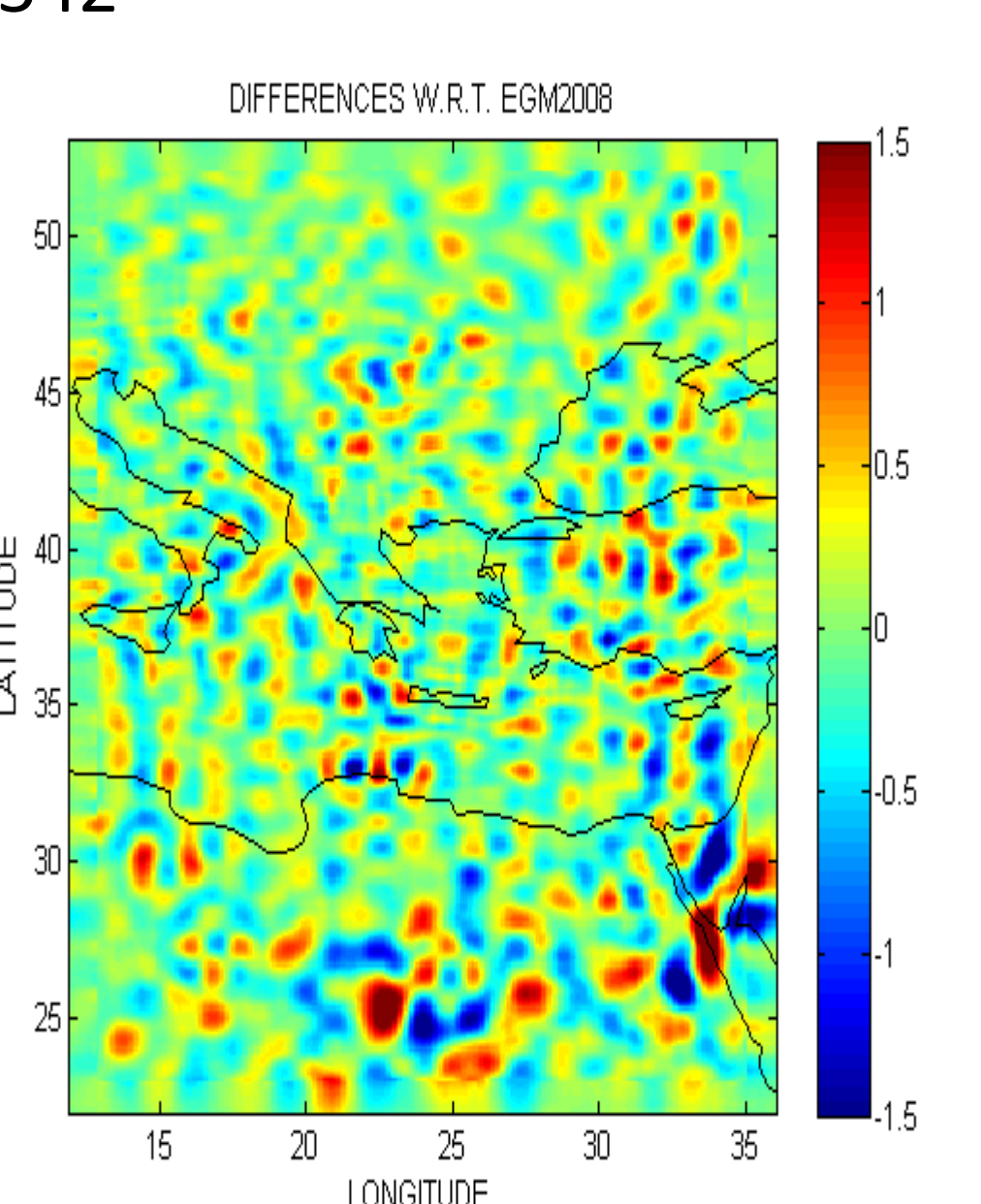


Figure 6: Combined geoid diff. w.r.t. EGM2008 [m]

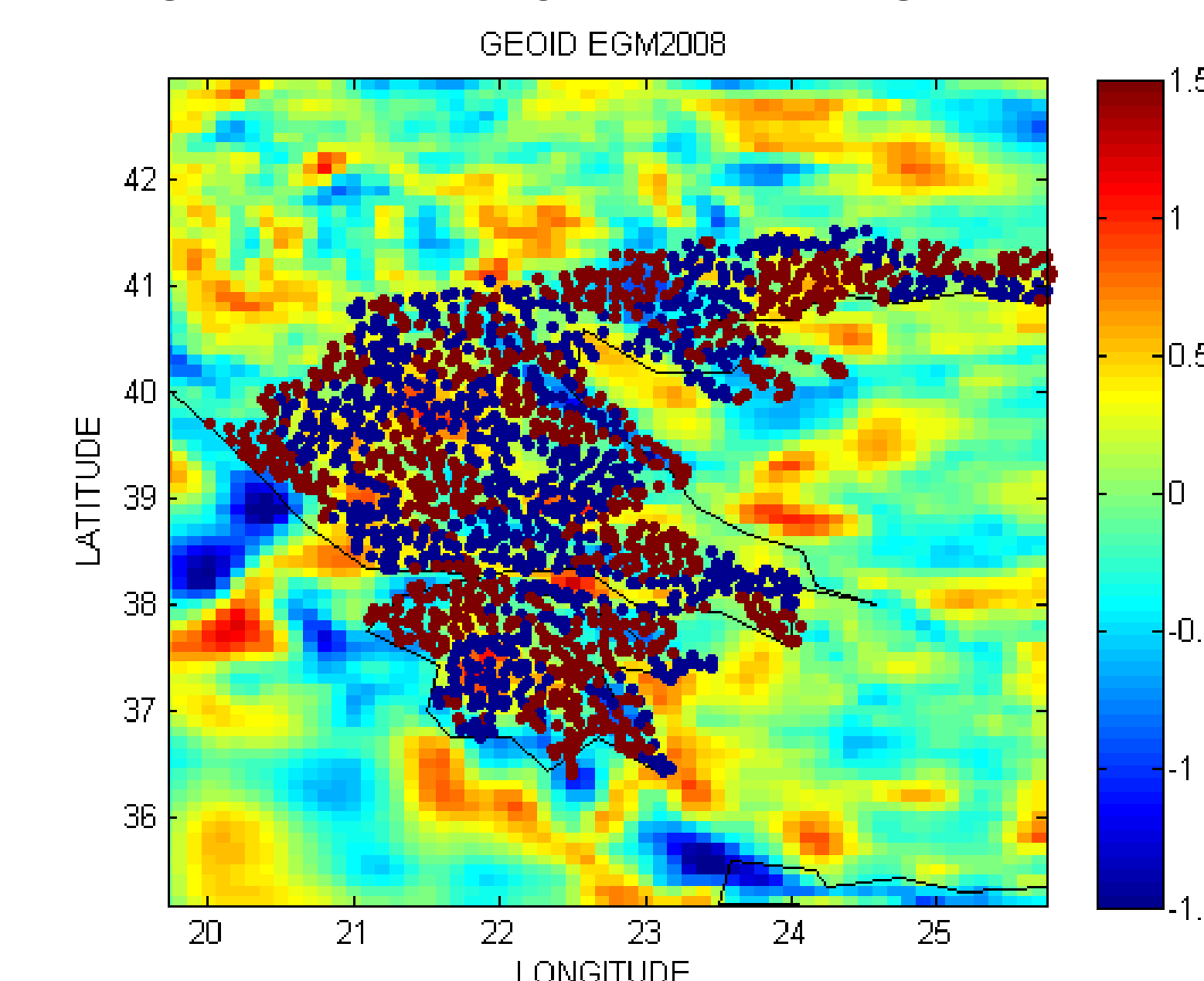


Figure 7: Combined geoid: improvement w.r.t. GPM DIR R2-blue: GPS/Lev BMs no improvement; brown: improved GPS/Lev BMs

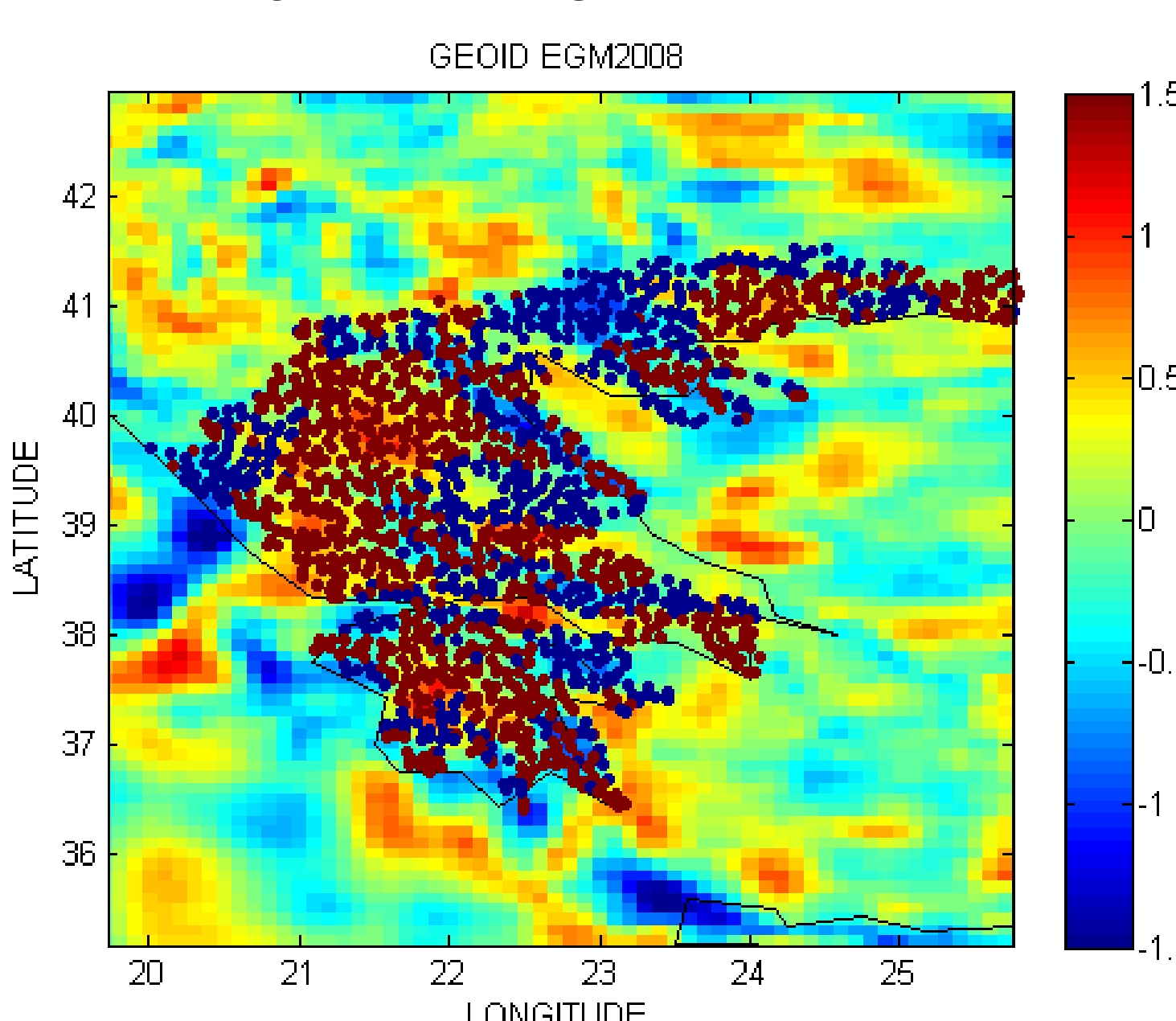


Figure 8: Combined geoid: improvement w.r.t. EGM 2008 blue: GPS/Lev no improvement; brown: improved GPS/Lev

	GPS/LEV	EGM	GOCE: Tzz	GOCE: Txx&Tyy	GOCE: COMBINED
MIN	-2.141	-1.118	-1.352	-1.230	-1.120
MAX	0.648	1.210	1.222	1.130	1.160
MEAN	-0.841	-0.011	-0.070	-0.029	-0.052
RMS	0.956	0.397	0.393	0.376	0.349
STD	0.455	0.397	0.387	0.376	0.345

Table 2: Statistics of geoid residuals at GPS/Lev BMs w.r.t. GPM DIR R4. Unit: [m]

	GPM	EGM	GOCE: Tzz	GOCE: Txx&Tyy	GOCE: COMBINED
MIN	-2.141	-2.553	-2.727	-2.775	-2.665
MAX	0.648	1.147	1.499	1.307	1.267
MEAN	-0.841	-0.831	-0.771	-0.812	-0.789
RMS	0.956	1.072	1.020	1.035	1.007
STD	0.455	0.678	0.668	0.641	0.627

Table 3: Statistics of GPS/Lev BMs' differences w.r.t. GPM DIR R4, EGM 2008, Tzz Geoid, Tzz,lp Geoid, Combined Geoid. Unit: [m]

GOCE Tzz: IMPROVED GEOID					GOCE Txx&Tyy: IMPROVED GEOID				
	w.r.t. EGM		w.r.t. GPM	w.r.t. EGM & GPM		w.r.t. EGM		w.r.t. GPM	w.r.t. EGM & GPM
MIN/#POINTS	-0.785	884	-0.921	801	-1.215	863	-1.050	739	497
MAX/STATUS	0.661	BETTER	1.352	BETTER	0.982	BETTER	1.160	BETTER	47.9%
MEAN/ IN %	-0.060	57.3%	-0.001	52.0%	-0.042	56.0%	0.018	47.9%	32.2%
RMS/#POINTS	0.257	658	0.354	741	0.320	679	0.345	803	1045
STD/STATUS	0.250	WORSE	0.354	WORSE	0.317	WORSE	0.344	WORSE	WORSE

Table 4: Statistics of improvement w.r.t. EGM 08, GPM DIR R4 and EGM&GPM

Conclusions

- The developed procedure for the downward continuation of GOCE SGG from mean orbit to the Earth surface using Simulated Annealing Monte Carlo method can be successfully applied for regional/local geoid modeling, providing more high frequency signal comparing to the GOCE only models (see figures 3-5).
- The residuals w.r.t. DIR R4 GPM (Table 2) at GPS/Lev BMs have much reduced mean value (~ 80 cm) compared to GPS/Lev derived geoid and decreasing STD (behavior similar to EGM); The combined geoid has the lowest STD and RMS while $T_{zz,lp}$ geoid has the smallest mean value.
- The residuals of GPS/Leveling geoid at BMs w.r.t. the other models (Table 3) have the smallest mean value and RMS, while the lowest STD is again for the combined geoid; Table 2 and Table 3 show the overall best performance for the combined geoid.
- In terms of improvement at GPS/Lev BMs (Table 4) the greatest improvement is w.r.t. EGM (for the combined geoid ~60%), and w.r.t. GPM DIR R4 (for T_{zz} geoid - ~50%).